Some Basic Concepts of Fluvial Geomorphology





Changes in valley morphology over "geologic time"

(Leopold, 1997, p 106)



Channel changes over years or decades

(Ward and Trimble, 2004, p 181)



Basic Concepts

- Force and resistance
- If exerted force is less than resistance, no change
- If exerted force is greater than resistance, there is a change to the channel or gradient
- Channel stability does not = rigidity
- Stable channels are in dynamic equilibrium; that is:
 - Channels don't change rapidly with time
 - The import and export of sediment is in balance
 - The channel is neither aggrading or degrading
 - Excessive mass wasting is not occurring

Morphological Characteristics

Channel shape is primarily a function of 8 variables (Leopold, Wolman and Miller, 1964):

channel slope, width, depth, discharge, velocity, roughness of the channel materials, sediment load, and sediment size.

Basic Stream Channel Shapes

(Rosgen, 1996, p 4-4)



Meanders



Sinuosity

- The ratio of channel to down valley distance
- Ranges from 1 to 4 in streams.
- Streams with a sinuosity of 1.5 or greater are called meandering streams.
- Streams with a sinuosity of less than 1.5 are called straight streams.
- Even in straight streams the thalweg tends to meander somewhat.

Meanders

- Water will follow the path of least resistance
- Pipe flow experiments have shown that bends that have a ratio of curvature to width of 2 exhibit minimum resistance to curvature.
- Meanders tend to follow a sine-generated curve shape:

Θ = ωsin (S/M) 2π

Where:

- **Θ** = the angle between direction and main valley direction at a given point
- ω = the maximum angle of deviation
- **S** = point distance along the path
- M = total path distance along a unit wavelength

Meanders (cont.)

Sine-generated curves represent the most uniform distribution of stress along a curve.





From: Leopold 1994, p. 66





⁽From: Leopold 1994, p. 62)

Weandering of a road ditch on deposited sediment (flat gradient).

Meander geometry

Repeated measurements of meanders across many stream sizes has shown a consistent relationship among the wavelength, width and radius of curvature:

wavelength is nearly always between 10 – 14 channel widths.

• wavelength averages 11 times channel width.

 the radius of curvature in the central part of the bend averages 1/5 of the wavelength.

 based on the relationships above, the radius of a meander is 2.3 times the channel width (1/5 x 11 channel widths = 2.3) The similarity of meander geometry across streams of different sizes is demonstrated by the 3 streams to the right that are drawn to the same size on the page.





Velocity Profiles in Meanders

Generalized surface streamlines

Generalized velocity distribution



Build up of a flood plain near meander point bars

(Leopold, 1997, p 79)













(From: Rosgen, 1996; after Schumm, 1968)

Straight Channels

- Straight channels have a sinuosity of less than 1.5
- Channel gradients tend to be close or equal to valley slopes.
- Stream energy is dissipated by the pool-riffle sequence. (and sometimes other structures)

The Pool and Riffle Sequence



Swiftcurrent River GNP, Montana

Cedar Creek, McCurtain County, OK



Colorado River, Cantonlands NP, Utah Q = 60,000 cfs

Pool and Riffle Geometry

- In a very general sense, the pool-riffle is a sequence of alternating deep and shallow water.
- Is another form of resistance to flow that dissipates stream energy.
- They form in gravel-bedded streams.
- In sand-bedded streams deep and shallow water and resistance to flow are provided by dunes and ripples.
- Riffle/pool spacing occurs at frequency of 5 to 7 bankfull widths.

Straight stream channels



Meandering or curved channels



How do pools and riffles form?

- The actual physical process is still unknown.
- In theory, as depth (discharge) increases, smaller particles should move first. If discharge continues to increase, larger particles will be transported (winnowing).
- Individual rocks interact with one another and influence their collective movement.
- Closely spaced rocks move downstream at a slower rate than those spaced further apart.
- They tend to bunch up in "platoons" which form riffles.



From: Leopold 1994, p 77

Bankfull and Effective Discharge

Bankfull and Effective Discharge

The concept of effective discharge:

- The effectiveness of erosion and deposition in forming stream channels increases with discharge.
- Low discharges occur frequently, but are ineffective in transporting sediment.
- High flows are effective in transporting sediment, but occur infrequently.
- An intermediate flow occurs often and is capable of transporting significant quantities of sediment.
- Effective discharge is the discharge that carries the largest amount of sediment over a long period of time.



Bankfull and Effective Discharge (cont.)

- Bankfull discharge is the flow that just fills the channel to the top of its banks and at a point where the water begins to overflow onto a flood plain.
- It occurs (exceeded) on average once every 1-2 years
- Or it has a return interval of 1.5 years



• Bankfull stage (or depth) is the incipient elevation on the bank where flooding begins.



Bluff Creek, Battiest, OK





10 year recurrence interval equalled or exceeded once every 10 years

Bankfull flow, recurrence interval 1.5 years, reached or exceeded 2 times in 3 years

Mean annual flow equalled or exceeded 30 percent of the time or 109 days per year

Low flow, expected 95 percent of time or 328 days in a year

Figure 8.3 The amount of water in a river channel and the frequency with which such an amount occurs.

For most streams, bankfull and effective discharge are the same



Watts Branch, Maryland

(From: Leopold 1994, p. 152)



Left Hand Creek, CO (From: Leopold, 1994, p 210)

Mountain Fork River, Oklahoma flow duration curves before (black triangles) and after (red circles) Beaver's Bend Dam



Baron Fork Suspended Sediment

